

DETAILED ALTERATION CORE-LOGGING, PETROGRAPHIC  
STUDY, AND A STRATIGRAPHIC-STRUCTURAL  
CONTROL OF MINERALIZATION HYPOTHESIS AT THE  
BROMIDE DISTRICT, NEW MEXICO-DDH B-1

APRIL 6, 1978

## SUMMARY

Two ore mineral assemblages were identified during the course of the study of DDH B-1, a sulfide assemblage made up of pyrite, pyrrhotite, chalcopyrite, and bornite, and an oxide assemblage composed of native copper and cuprite. Generally, the sulfide assemblage was localized concordant with foliation and in association with carbonate-hydrothermal chlorite stringers. The localization of the oxide assemblage is along fractures and foliation surfaces where groundwater could penetrate.

An attempt to correlate rock-type and mineralization showed that higher copper values were associated with agglomeratic, felsic units. These units may or may not have been the primary carrier of values, but subsequent metal mobilization may have localized values here due to the more porous nature of the rock.

Generally, chlorite, sericite, biotite and epidote are the most important alteration minerals. Microscopic studies revealed the widespread character of sericite, the replacement of early chlorite and biotite of metamorphic origin by latter hydrothermal chlorite, and epidote as a replacement product of plagioclase and not uncommonly disseminated in the deeper portions of the drill hole. Other alteration phases observed were serpentine, stilphomelane, and cordierite. The juxtaposition of an altered unit over an unaltered unit, indicating possible inverted stratigraphy, suggests the identification of a marker unit to identify the top of a single mineralization episode.

By plotting top and bottom indicators down the drill hole, cartooning overturned isoclinal folds, and superimposing anomalous copper values on the picture, there is suggested a possible structural-stratigraphic control for mineralization. Altogether, the evidence suggests further search either up-dip in the same stratigraphic unit, or down strike.

## ORE MINERAL ASSOCIATIONS

Whether the ore components were present at the time of deposition of the alternating sequences of andesites, agglomerates, sediments and ashflows during the Precambrian or not is of academic importance with respect to their observable, present-day distribution. Generally speaking, mobilization and re-deposition in favorable localities characterizes the sulfide-oxide ore mineral assemblage, resulting from both hydrothermal and supergene mechanisms.

Sulfide Assemblage. Pyrite, pyrrhotite, chalcopyrite and bornite make up the assemblage of ore minerals deposited as a result of hydrothermal processes active late in the greenschist metamorphic episode or following it. The zones of weakness (foliation) which localized the sulfide minerals pre-date their deposition, indicating a relative age for mineralization. The ore minerals are very commonly in association with quartz-carbonate stringers which rarely cross-cut the pervasive structural grain of the host rock. In instances of obvious post-metamorphism shearing, accompanied by subsequent quartz-carbonate healing, remobilization of the copper and iron sulfides has not occurred, leaving the fault zones quite barren.

The typical association of ore and host materials consists of any single sulfide phase intergrown with, or sheathed by a secondary, bright green variety of chlorite, commonly associated with calcite. In rare instances, chalcopyrite and bornite occur together, while on occasion pyrrhotite was observed to replace magnetite. Generally, pyrite is the most common ore mineral and occurs in the upper portions of the hole, while pyrrhotite, possibly resulting from late de-sulfurization of pyrite, predominates in the bottom 300 feet.

Oxide Assemblage. Native copper and cuprite make up the supergene assemblage of ore minerals. Controls for the distribution of these phases are most commonly planes of foliation, and sympathetic fracture surfaces which acted as conduits for meteoric water in the subsurface. Native copper occurs as thin plates on limonite-coated surfaces. Cuprite is more closely associated with blebs and veinlets of silica, and is restricted to the bottom half of the hole.

#### CORRELATION OF ROCK TYPE AND MINERALIZATION

The consideration of ore grade as it is related to lithology is an aspect that should not be overlooked since an obvious guide to mineralization may become evident through such correlation. Fifteen assay samples from DDH B-1 whose copper values were in excess of 250 ppm were noted along with footages associated with each sample. Matching the assay interval with the alteration log interval yields an interesting corelation of rock and assay value, and implies a genetic relationship as well as suggesting a source of heat to mobilize and redeposit copper, chlorite and carbonate of hydrothermal affinity. In addition, implication of location laterally displaced from higher grade zones may be put forth.

A striking correlation of higher mineral values with agglomeratic, fragmental, or brecciated materials is immediately obvious from the following table. It may be suggested that these units not only carried initially anomalous concentrations of copper (syngenetic), but also that the more porous character of a rock of this sort would be more receptive to the migration and re-deposition of mobilized primary copper. Perhaps it may not be too far out of line to suggest these materials as being "millrock" equivalents.

A noteworthy interval from 350 feet to 510 feet exhibits anomalous copper over its entire length. It is suggested as one possibility that heat from a nearby intrusive rock, evidenced by observed, small apophysés of later trap material into this sequence, may have been responsible for localizing copper values and imparted the hydrothermal character to the distribution of chlorite, carbonite, and sulfides. A second, equally plausible explanation suggests that these units are laterally contiguous with, or equivalent to higher grade units of an ore-depositing episode. Therefore, exploration drilling stepped out along the strike of these beds may show greater promise.

#### MEGASCOPIC PHYLLOSILICATE DISTRIBUTION

From megascopic observation of split core from DDH B-1 it was observed that these phyllosilicate phases predominate: sericite,

chlorite, and biotite. The distribution of these minerals appears to be governed by two controls, the rock type and the alteration/metamorphism intensity.

The initial 330 feet down-hole is dominated by sericite, with subordinate amounts of chlorite. It is believed that the upper 75 to 100 feet of the section is characterized by supergene sericite resulting from surficial weathering effects. Deeper into the hole, significant amounts of sericite mixed with chlorite occur within all rock types. Narrow intervals, usually of meta-andesitic composition, in which chlorite is the dominant sheet silicate, occur at irregular intervals down-the-hole.

Between 330 and 660 feet chlorite is the chief mineral resulting from metamorphism/hydrothermal alteration. Minor amounts of sericite, except in a sericitized meta-ashflow unit, together with the initial appearance of biotite, are also part of the mineral assemblage. Chlorite of metamorphic origin is easily discernible from that originating by hydrothermal processes. The metamorphic type is generally pervasive in its distribution, and is very dark green to blackish-green in color. Most commonly it is quite fine grained. In contrast, later hydrothermal chlorite is bright lime green and coarser grained. The distribution of this generation of chlorite is usually controlled by veinlets in association with quartz, carbonate and occasionally sulfides. Not uncommonly, it is observed as thin laminae concordant with the foliation of the metamorphic rock. In zones where concentrations of this secondary chlorite are high, the probability for ore materials is greatest.

Biotite and sericite  $\pm$  chlorite dominate the lowest portion of DDH B-1 (660 feet to 830 feet). Biotite is usually localized in laminae and near veinlets as medium to fine-grained books aligned parallel to foliation. Its presence indicates increased metamorphic intensity, therefore higher P-T conditions. Winkler (1967) suggests that coexisting muscovite (sericite) and chlorite react to form biotite and aluminum-rich chlorite or muscovite. The light grey, platy mineral common in the deeper sections of the drill core is possibly

TABLE 1. INTERVALS WITH  
COPPER ASSAYS > 250 ppm

<u>SAMPLE NO.</u>	<u>ASSAY INTERVAL</u>	<u>LOG INTERVAL (S)</u>	<u>CU</u>	<u>ROCK TYPE</u>
B-1-2	77-97	71.5-81 81-88.5 88.5-97.5	270	Agglomerate-andesite porphyry
B-1-8	158-168	155-165	265	agglomerate
B-1-29	350-360	348.5-356 356-374.5	260	andesite-sediments
B-1-32	380-390	379.7-385	275	andesite-agglomerate
B-1-33	390-400	385-398.5	265	andesite
B-1-37	430-440	398.5-439	275	andesite with pebbles near intrusive
B-1-39	440-456	439-473.5	810	" " "
B-1-40	456-470		315	" " "
B-1-44	500-510	494.5-512.5	420	agglomerate-qtz vein
B-1-49	546.5-560	547-559.9	260	ashflow
B-1-51	570-580	559.9-579.5	305	ashflow with freqs
B-1-52	580-590	579.5-592	255	andesite
B-1-56	620-629	614-625 625-629.3	285	andesite
B-1-57	629-633	629.3-641	305	agglomerate
B-1-75	788-789	770.2-803	6000	andesite with auto- liths

this latter reaction product and accounts for the low chlorite observed once biotite appeared.

#### MICROSCOPIC-PETROGRAPHIC INVESTIGATION

Among the important results of the microscopic study of selected sample intervals from DDH B-1 drill core were the paragenetic interpretation and alteration phase distributions, and the identification of a marker unit that may signify the top of a metallogenic episode in the volcanism.

In terms of paragenesis, replacement textures in which phases interpreted to be hydrothermal replacing the "primary" metamorphic minerals are deemed most important, since it is this hydrothermal event that is thought to be associated with the ore mineral mobilization and relocation. Starting with rocks initially containing quartz, plagioclase and mafic minerals, usually fineequigranular in character, an alteration assemblage composed primarily of sericite, chlorite, epidote and calcite, with minor amounts of serpentine, stilpnomelane and cordiente, was superimposed on a chlorite, plagioclase, biotite, sericite metamorphic assemblage.

Sericite was observed to be much more widespread than originally perceived megascopically, and was noted to replace plagioclase with varying degrees of pervasiveness, cordierite to a minor degree and chlorite-1 (metamorphic origin) in small amounts. Chlorite-2 (hydrothermal origin) was observed to replace biotite, chlorite-1, and in a moderate manner, plagioclase. Additionally, veinlets and quartz-carbonate pods tended to localize chlorite-2. Epidote was sparse near the top of the drill hole, but the extent of its distribution increased significantly below 550 feet to a maximum of 20-25% of the entire rock. Generally, epidote was restricted to plagioclase replacement, but occasionally it was disseminated as small grains throughout the rock, locally concentrated along inter-unit contacts. Calcite was very common as a veinlet mineral in association with quartz and chlorite-2. However, it was not uncommon to observe calcite to replace plagioclase phenocrysts in a splotchy manner.

Serpentine was usually observed in groundmass location and probably was formed as a result of replacement of earlier mafic grains. Cordierite, exhibiting its characteristic interpenetration twinning, was rarely observed.

Stilpnomelane, a member of the brittle mica family, was observed along several large calcite veinlets and probably was generated by the late hydrothermal activity.

Metamorphic biotite was first noted at 543 feet, indicating increased metamorphic intensity at great depths. Coarsely crystalline sericite, megascopically interpreted as argillization, was noted at 462 feet. This corresponds quite well with the high copper assay values. It is suggested this may be a recrystallization product associated with the hydrothermal event.

Near the bottom of the drill hole a lightly altered andesite porphyry was observed to stratigraphically underlie a highly altered sequence. This would again suggest the overturning of units with mineralization being stratigraphically lower in the sequence. This andesite has the potential to be a marker unit since it has discernible plagioclase phenocrysts replaced by epidote set in a very fresh, all plagioclase groundmass. If there has been only limited sulfide mobilization, this unit possibly marking the top of a mineralization episode, will be useful.

#### BEDDING ORIENTATION & POSSIBLE INTERPRETATION

Study of the drill core from DDH B-1 revealed conflicting evidences for the determination of tops and bottoms of the defined rock units. As a result, the author suggests the possibility of folding within the sequence. By cartooning the evidence observed in the drill core for bedding orientation, the following figure was derived. By superimposing the anomalous Cu values from Table 1 at the appropriate footages, an interesting effect became noticeable. The high copper values ( $>250$  ppm) seemed to correlate well with the overturned limbs of the hypothetical folds, each time quite close to a proposed crest.

The "tightness" of the folds was inferred by the fairly constant dip observed in the bedding and the frequency at which top and bottom indicators were noted, suggesting an isoclinal folding pattern. It also seems worthwhile to note that the copper anomaly seems to be restricted to a particular unit.

If this hypothesized stratigraphic-structured relationship in any way reflects the natural situation, the suggestion for ore localization might be as pods or rod-like bodies within agglomeratic felsic units near fold crests. The correlation of values from the Payroll Shaft seems to corroborate this model quite well. Further investigation up-dip toward the hypothetical fold crest and along the strike of the units seems to be merited.

Robert W. Schaefer

DDH B-1  
(-55° N 45° E)

PAYROLL  
SHAFT  
(PROJECTED)

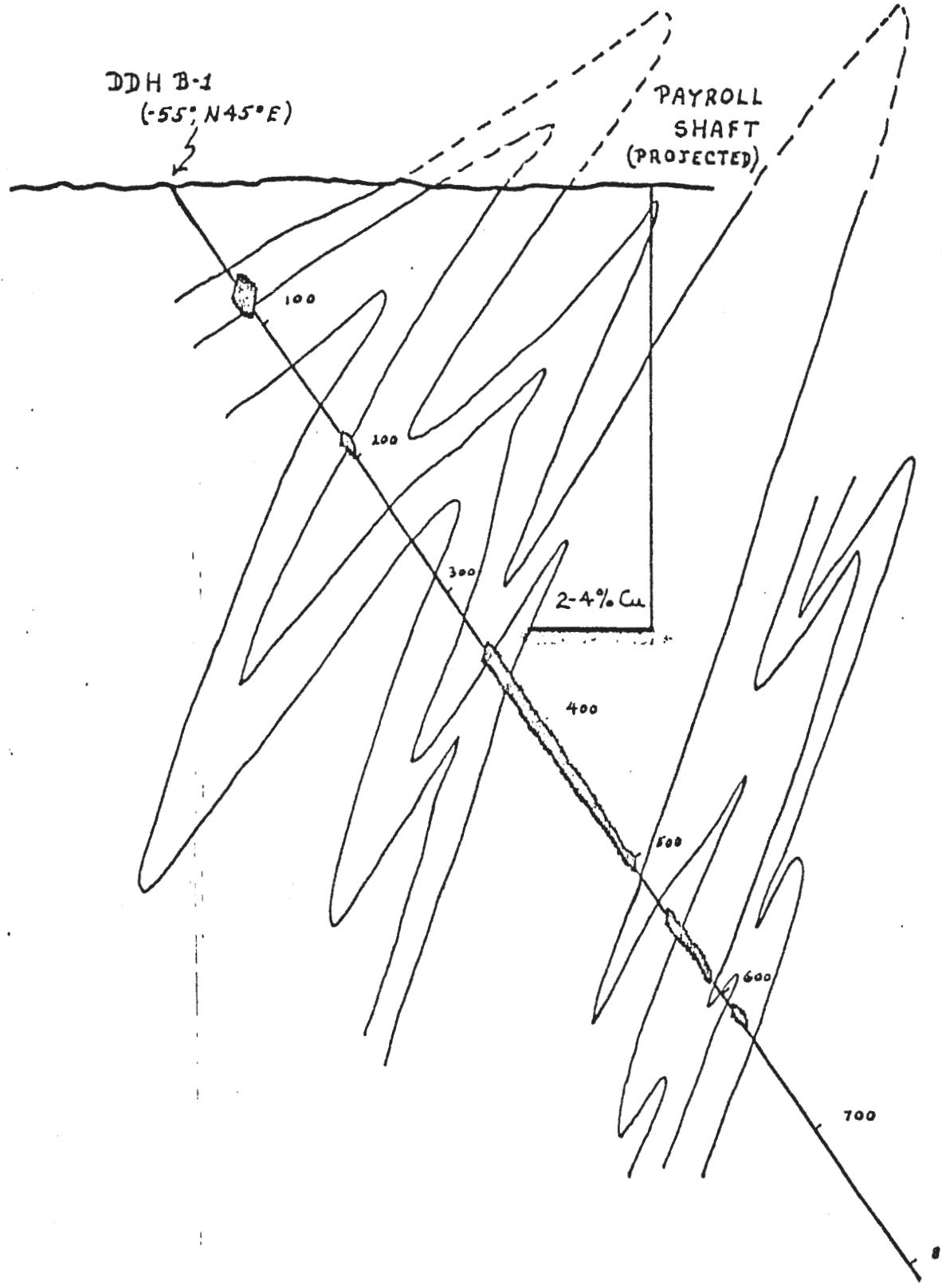
Cu ASSAY >250 ppm

HYPOTHETICAL SECTION  
THROUGH DDH B-1  
(LOOKING N 45° W)

BROMIDE PROJECT  
RIO ARRIBA COUNTY,  
NEW MEXICO

SCALE 1 inch = 100 feet

7 APRIL 1978 R.W.S.



USBORAX

DIAMOND DRILL LOG  
EXPLORATION

## HOLE SURVEY Detailed Alteration.

Footage

Bearing

Inclination

LOG

Property BROMIDE PROJECTHole No. B-1Location Rio Arriba Co., New Mex

Total Depth \_\_\_\_\_

Date Started \_\_\_\_\_

Collar Elevation \_\_\_\_\_

Date Completed \_\_\_\_\_

Incl. &amp; Bearing \_\_\_\_\_

Logged By R.W. SCHAFER

Core Size \_\_\_\_\_

Stage	Int.	Rec.	Alteration*				Mineralization			Graphic Log		General Description Rock Type, Structure, Etc.		Assays			
			% Chl	ser	cal		T.S.	Ore	Min.								
0	10	10								Fe-oxides			SURFACE PIPE (0-70 CORE BOARD)				
0	20	10	5	60							10		60% SERICITE - 25-30% QUARTZ - 5% CHLORITE SCHIST - ACCESSORY MAGNETITE (2-4%) - LIMONITE ALONG SCHISTOCITY (HEMATITE-RICH) after MAGNETITE PYRITE - MODERATELY OXIDIZED - META-ANDESITE -				
0	30	10	5	60							20		SAME AS ABOVE - SERICITE IS LT. GREEN TO WHITE - LIGHTLY OXIDIZED - POSSIBLY SERICITE IS SUPERGENE in ORIGIN				
0	40	10	30	30							30		CHLORITE-SERICITE-QUARTZ SCHIST - CHLORITE & SERICITE ABOUT EQUAL - LIMONITE PSEUDOMORPHS after PYRITE (HEMATITE-RICH) - MAGNETITE REMAINS ACCESSORY meta-andesite				
0	50	10	30	30							40		SAME AS ABOVE - LIMONITE PICKING UP A LITTLE more JAROSITE (HEM 85, JAR 10%) - COARSER, BULL-EYE FRAGMENTS (VEIN?)				
											50						

Footage	Int. Rec.	Alteration *					Mineralization			Graphic Log		General Description			Assays				
		From	To	%	Chl	Ser	cal	Flr		T.S.	Ore	Min.	Gangue	Ft.	Rock Type, Structure, Etc.				
50	60	10			45	15							Fe-oxides		CHLORITE-SERICITE-Quartz SCHIST - Limonite Localized along SCHISTOCITY (jarosite rich) - hematitic Limonite after mag, PY (?) - higher Chlorite - magnetite is EUDERAL to subhedral octahedra				
60	70	10			45	15								60	SAME AS ABOVE - LESS OXIDIZED				
70	71.5				20	45	1-E	TR		Diss. PY-Fe-TiO <sub>2</sub>	MAG → Z-51%			70	C-1-02 STRANGERS, OFTEN LOCALIZING PY; SLIGHTLY ROTATED PLAG PHENOS-CATACLASTIC rim healed by QTZ, TRACE FLUORITE (?) in. CAVITY..... V. fine grained SER-CHLOR-gtz phyllite	THIN SECTION	B-1-71	B-1-75	B-1-83
71.5	81	9.5			55	15	4-6			tr. PY	MAG → 8-9%				CHLORITE STRANGERS w/ ASSOC magnetite = SUBHEDRAL CUB. MAG ASSOC w/ CHLOR STRANGERS coarser than diss. mag. vein controlled Lim-Hem, gtz lenses, few relict plagiophenites - meta-andesite TO meta-agglomerate gtz veinlets often have mag dust HALO - CLEAR pc QTZ-calcite, DH's calcite - INVERTED GRADED PEBBING at 78', following A SUETLE TRUNCATION - gtz-chlor-mag pseudomorph after biotite (?) - PERLICLES OF Agglomeratic intervals stratified along foliation direction				
81	88.5	7.5			40	30	2			LIM along SCHISTOCITY MAG DUST in CHLOR MAG → 6-8%				80	at 78', more oxidized mag → hem common and higher Sericite content (~30%) ASSOC with agglomeratic interval TO 81'				
88.5	97.5	9			5-10	60	4			coarser mag, diss + veinlets 2-4% up to 7-8% - trace ilmenite with leucocore 1-2% AT 90°				90	AGGLOMERATIC UNIT, DOMINANTLY SERICITE, CALCAREOUS LENSES, ROTATED GTZ PHENOS, SOME WITH CATACLASTIC GTZ PHENOS LACK SHARP BOUNDARIES, RATHER BLENDING INTO QTZ-SER-CHLOR MATRIX, - CALCITE STRANGERS, LENSES & DISSEMINATED - MUCH MORE GTZ-RICH WITH DEPTH (TOP AFUAR?) SAME WITH SER, YELLOW CALCITE COMMONLY w/ QTZ	THIN SECTION	B-1-75	B-1-83	B-1-89
97.5	108	11.5			15-20	55				magnetite DUST AMB & FEW COARSER XCRALS 3-5%				100	Dominantly, gtz-ser-mag UNIT, CONTACT RELATIONSHIP INDICATES PROBABLY EARLIER THAN PRECEDING INTERVAL, SOME LENTICULAR CHLORITIC ZONES				
108	123	15			30	35	16			ACC. magnetite 2-3%				110	SER-CHLOR-QUARTZ PHYLLITE POSSIBLE LT. green FLUORITE AS LATE OPEN SPACE FILLER, CONCORDANT GTZ LENSES AND MICROSTRANGERS, OFTEN LIMONITE STAINED A FEW GTZ EYES, ALSO VEINLETS OF LIMONITE-HEMATITE-GTZ A FEW REFLICT "GHOSTS" OF HORNBLENDITE (?) PHENOS REPLACED BY VERY FINE CHLORITE AND MAGNETITE META-AGGLOMERATE - MOTTLED, STRETCHED PERLICLES COMPOSED DOMINANTLY OF SERICITE AND QUARTZ SURROUNDED BY CHLOR-SER-GTZ MATRIX, VARIABLE DISTRIBUTION OF CALCO-SERICITE SOME GTZ VEINLETS WITH LIMONITE HALO'S (AFTER MAG?) INFREQUENT CLOWS OF CHLOR-WRTE (AFTER PLAGIO?)	THIN SECTION			

Page	Int.	Rec.	Alteration *				Mineralization			Geologic Log		General Description	Assays	
			% CHL	Ser	Cal	FLR	EP	T.S.	Ore	Mn.	Gangue	Ft.		
INT TO														
3	141	18	60 TO 45	15'	6-7						2-3% mag lightly oxidized on rims and commonly pseudomorphically replaced by hematite-martite (disseminated)	130	FLUORITE IS LATE STAGE, TEXTURE BECOMES FINER WITH DEPTH - LOSES AGGLOMERATE CHARACTER AND BECOMES MORE EQUIGRANULAR, FLOW-LIKE RUGY	
													CHLORITE - QUARTZ - SERICITE PHYLLOLITE - Homogeneous texture - EQUIGRANULAR limonite stained quartz and disseminated yellow calcite and clots that cross-cut semistability - Disseminated calcite may be replacing plagi meta-andesite	THIN SECTION B-1-12
H	145	4	20 TO-11	10	8						HEM. after mag-martite 1-2%	140		
J	149	3.5	20	30-35	5-6						mag oxy, larger biss grains jarosite limonite	150	meta-gtz diorite-andesite, STRETCHED gtnos mafic (hornblende?) replaced by CHLORITE	
K	150.5	1.3	60-40	5-20	5-7						1-2% mag ..	150	- CHILLED margin, pheno population increases away from contact (WITH DEPTH), BECOMING EQUIGRANULAR gtz-cal veinlets	
L	153	2.7	20-15	35-60	2-3						mag, hem 3-4%	150	calcite also in PLAG sites with gtz and relief plagi	
M	155.5	2.5	60-55	5-10	1-2						TR PYRITE MFORE W/ ORE	Py oxidized TO limonite	- lightly oxidized, becomes relatively EQUIGRANULAR and more stretched at depth - DIKE (?)	
													- A few XENOLITES near contact at depth (maybe formerly a dike?)	
													CATACLASIZED gtz-CHLORITE VEIN, HEATED BY CALCITE	
													meta-sedimentary texture, veinlet gtz	
													CHLORITE RIM around stretched pebbles	
													GTZ-CHLOR-SER ARGILLITE w/ acc magnetite	
													magnetite incr w/ depth (heavies settling out, right side section)	
													- gzt 148+ or Cu, oxidized PYRITE	
													- CHLORITE occurs as conformatable strangers and replacements for stretched grains and running gtz, high Limonite content at contact	
													- GRADED bedding, COARSER frags w/ depth	
													- next unit has ONTREIVE, OR later emplacement DIAKE - veinlets include iron textures	
													CHILLED margin, v. fine grained CHLORITE, BECOMES POEOPHYITIC TOWARD INTERIOR - REPLACED BY GTZ-CAL,	
													more SERICITE TOWARD CENTER	
													- GTZ-CAL VEN - DIAKE -	
													GTZ-SERICITE-PLAG MAGNETITE-CHLORITE	
													MELA-SED ROCK, BECOMES MORE SERICITE & MAGNETITE RICH w/ DEPTH, CORRELATES w/ 145-149 INTERVAL	
													Chlorite-diorite-melange, becomes "poephylitic" in final 15' w/ cal. and lamellae continuing to foliation more w/ depth w/ diake, stretched CHLOR PLEOBOIMPHS after metam - normal - GRADED bedding?	THIN SECTION B-1-153





Footage from To	Int. No.	Alteration *					Mineralization			Graph	Log	General Description	Assays	
		% CHL	SER	CAL	FIR	EP	T.S.	Ore	Min.					
67	269	Z	30	20	TR-1					mag DUST, ODS BENNS 1-2%		CHLOR-SER-CALITE PHYLLOLITE - v. fine grained, slightly foliated, calcite stringers w/ CHLOR- mag haloes - CP assoc.	THIN SECTION B-1-268	
69	311	40	10	40	TR-1					mag DUST assoc w/ CHLOR in stringers	270	Qtz-Ser-mag SCHIST w/ CHLOR-CAL stringers, quite fine grained, mag in lenses, rare & CRYST CYSTS; TRACE CAL is rock, most in stringers; Rare stretched frags near top of unit, grading down to a pebbly unit (agglomerate) with stringer CHLOR-(CAL-CP) between SER-CP frags w/ mag dust; CHLOR CLOTS in QTZ-SER probably as major replacement, in addition to CHLOR. Stringers - CHLOR is coarse & outlined in QTZ- CAL-CHLOR veins; Becomes evenly grained dominantly QTZ-SER w/ CHLOR-CAL stringers - SULFIDES assoc with CHLOR exclusively, local more schistose units (2-6" thick) are higher in CHLOR (25-30%), but generally remains QTZ-SER and as low in SULFIDE content, general incr in carbonate stringers from 287 to 290.5, and from 292.5 to 298.5'. - CAL-CHLOR stringer w/ oxidation halo - no mag general dark color due to DISS mag & fine grained nature of minerals, zone of more intense CP lens at 290'; a few small green eyes; SCHISTOSITY increases in ZnO zone or cal stringers (due to high CHLOR content?), also a few stretched qtz pebbles - with CHLOR-SER in fractures, pebble population increases with depth, ROTATED somewhat DISRUPTING FOLIATION, CAL-CHLOR- RELICT PLG (?) - QTZ CLSTS, some SERPENTINE along fracture cross-cutting foliation (0.8'), more apparent foliation from 302 to 305 with greater SER:CHLOR Some pyrite w/ minor cal; a few small, stretched pebbles, lge frag at 306 setting in v. fine grained QTZ-CHLOR PHYLLOLITE, frag is compositionally similar to interval 298, contains CP - surrounding rock does not, CHLOR HALO becomes a QTZ-CHLOR- SER phyllite w/ a few pyrite lenses, partly replaced by Calcite, CHLOR content incr, as does schistosity	THIN SECTION B-1-298	
1	332	21	35	TR	10					TR PY, CP	300	FRESH mag - a few grns, mostly DUST	Shear contact, slight drag on schistosity; CHLOR-CAL-mag SCHIST, cal replacing plg ad as veins; Becomes finer v/ depth, more EQUIGRANULAR w/ SERPENTINE becoming more noticeable; short interval of fractured, finely grained CHLOR- SER-ATZ PHYLLOLITE's w/ CHLOR healing fractures, then goes back to a fine gr. CAL-CHLOR phyllo - this pattern of CHLOR-CAL ad CHLOR-SER-CAL phyllites alternates w/ every 2 ft. CHLOR-CAL stringers in serp-qtz-CHLOR rocks, Local Haloes around CHLOR- CAL-mag pseudomorphs, some stretched QTZ-SER concentric lam at 325' - maybe after plg or pebbles	
		AN6	35	25	4						310			
			↓	↓							320			
			30	10-15							330			









Footage from To	Int. Sec.	Alteration *					Mineralization			Graph	Log	General Description		Assays
		% CHL	Ser	cal	FLR	EP	T.S.	Ore Min.	Gangue			Ft.	Rock Type, Structure, Etc.	
285	530.5	2	65	5	5			TR. cuprite	NO mag		520		BRECCIA UNIT - large clusters of unrecrystallized unit together w/ crystals of v. fine grained massive CHLOR UNIT with a CHLOR matrix - maybe some TALC also present; Qtz-cal assoc with fractures; CLASTS are fairly well-rounded	
302	536.0	5.5	45	30	2			TR Py	Hematite				medium grained CHLOR-Qtz-cal pyrophyllite meta-volcanic (andradite?) with LT gray sericitic & some PHENOS; Bright green CHLOR replaces mafics, Relict cal-carb & Qtz-Ser replaced phg and TR, a little clayey (?); faint Contact separating a lower CLASTIC poor metavolcanic from overlying CLASTIC-rich metavolcanic unit (diffuse flow?)	
36	542.0	6.8	35	15	3			—	Hematite & magnetite		540			
428	547	4.2	20 ↓ 30	TR ↓ 25	12		TR		—		540		May be metased or meta-volcanic CLASTIC UNIT; Similar to Cal Unit, green CHLOR-LT gray Sericitic, Qtz-Calcite; rounded, slightly elongated Qtz-Cal pebbles, some evidence for slight rotation of pebbles; small lenses of mag, CHLOR, Qtz eyes become common near lower contact of unit suggesting more of sed origin (meta-graywacke?), may rich new base	
47	559.9	12.9	75 ↓ 20	35 ↓ 55	TR ↓ 2-3		TR	TR Py at 558, trace Ba	Cubic mag 4%		550		Very fine grained CHLOR-Qtz-PLAG-mag-CAL "Hornfels" very little evidence for long range preferred orientation of grains; Qtz ~45%, PLAG ~25% near upper contact; w/ ser in CHLOR increasing near lower contact; may be some biotite (need T.S.), a thin meta-albite unit near upper contact	THIN SECTION B-1-54
59.9	579.5	19.6	65 ↓ 35	TR ↓ TR mag?	8		TR	TR CuO?	MAG CUBES DIST ~6%		560		fine grained Ser-CHLOR pyrophyllite, NO Calcite, Small rounded CLASTS; all is replaced by CHLOR-ser, the CLASTS are slightly recrystallized in CHLOR, Some lenses appear to be serpentine calcite only in plagi sites, a few sm. Qtz eyes, maybe a meta-ashflow + slightly streaked D "pumiceous" frags, alabard sugary Qtz; at 557, a thin andesite flow unit ~6" thick (~70% CHLOR, trace, 5% Cal), Numerous Qtz-Ser, rounded frags - may be recrystallized mafic-spherulites	THIN SECTION B-1-559
													Bracchia finger w/ Bracchia matrix at ~550, CHLORITE pseudomorph heavily replacing pyroxene, some CHLOR replacing mag	
													Contact w/ overlying unit shows that upper unit is earlier. See down sequence; Rock unit is laterally meta-andesite giving way to another Ash-flow. Dashed margin on Mafic unit, also a south of Ash-flow in andesite, CHLOR is bright green, TR, Epitrope after PLG, w/ col, Qtz, andesite is ~6" thick giving way to Mn. Relict PLG, pores of alumite? gypsum? MnCO <sub>3</sub> w/ cu in veins	THIN SECTION B-1-560















## BROMIDE PROJECT DDH-B-2

### SUMMARY:

Drillhole B-2 of the Bromide, N.M. project penetrates metavolcanics, metasediments and intrusive igneous phases similar to those of DDH B-1. However, the rock units intersected generally have been metamorphosed to a slightly greater degree than those observed in the previous drillhole. The greater abundance of biotite and epidote in the metamorphic mineral assemblages near the surface, along with the possible occurrence of garnet at greater depths are evidence of this effect.

The sulfide mineral assemblage is similar to that of DDH B-1 (pyrrhotite, pyrite, chalcopyrite, bornite) with the exception that a new mineral, possibly niccolite ( $NiAs$ ), was observed in association with pyrrhotite near the bottom of the hole. In addition, frambooidal pyrrhotite was observed on several fractures and foliation surfaces, suggesting biogenic activity following the metamorphic event(s). The oxide ore mineral assemblage consists almost exclusively of native copper; cuprite was observed only once.

New structural features were observed to corroborate suggestions made in the report for DDH-B-1. Overturned isoclinal drag folds were noted in shallow portions of the drill core to support the regional structural picture constructed from evidences of overturning in DDH B-1 core. Another fracture seen in sawed core sections was a lineation of the platy minerals that is not only at an angle to the primary bedding foliation, but also varies in its orientation down the hole. It has been interpreted from those features that another system of folds may be superimposed on the earlier set, not distorting the first, but rather occurring as a response to the closer proximity of these rock units to an intrusive body.

### ROCK UNITS

A variety of metamorphosed volcanic, volcanoclastic and sedimentary rocks are observable down DDH B-2, along with later unmetamorphosed igneous intrusive units. The greatest volume of rock is composed

of meta-agglomeratic units and andesitic units. Porphyritic flow phases and intrusive phases are also present, and those such as ones occurring at 355-363 feet or 395-400 feet appear to correlate to similar rocks in DDH B-1. Intervals from 390-395 feet and 525-560 feet are fragmental in character, of felsic nature, and may be rock units termed "millrock."

Other rock types making up significant, but lesser portions of the section in this area are ash flows and greywackes, both metamorphosed, as well as later dikes (Lamprophres) and fresh biotite diorite intrusive at 1113-1131 feet.

#### SILICATE MINERAL ASSEMBLAGES

The mineral assemblages which dominate the length of DDH B-2 are: 1.) Biotite-chlorite-sericite, 2.) biotite-sericite-epidote, and 3.) biotite-chlorite-epidote, although short intervals in which any two of chlorite, sericite, epidote, or biotite co-exist exclusively were noted. One interval exceptionally rich in chlorite resulting from the retrograde metamorphism of biotite has been recorded at 930-1040 feet. Between 1180 and 1185 feet a very fine-grained metamorphic mineral resembling garnet was observed. Petrographic study would be needed to confirm the identity of this red mineral.

Table 1 lists the dominant minerals present in various intervals down the hole. Greater description of the mineral relationships are available in the DDH B-1 core-logging report.

#### ORE MINERAL ASSEMBLAGES

Sulfide. Pyrite, pyrrhotite, bornite, and chalcopyrite are the most commonly observed sulfide minerals seen in the B-2 drill core. Generally, pyrrhotite and bornite are the most common of the group and usually co-exist. The fact that those, rather than pyrite-chalcopyrite as in B-1, are the most common may be due to the higher metamorphic grade developed in the B-2 core.

TABLE 1.

<u>Footage</u>	<u>Mineral Assemblage</u>
66-124	Chlorite-sericite-epidote
124-288	Biotite-sericite-chlorite
288-338	Chlorite-sericite-epidote
338-349	Chlorite-sericite
349-365	Sericite-biotite-epidote-(chlorite)
365-471	Chlorite-sericite-biotite
471-496	Chlorite-sericite-biotite
496-530	Chlorite-sericite-biotite
530-612	Sericite-biotite-epidote
612-621	Biotite-chlorite-epidote
621-693	Biotite-sericite-biotite-epidote
693-705	Chlorite-sericite-biotite-epidote
705-726	Biotite-sericite-epidote
726-730	Biotite-chlorite-epidote
730-855	Biotite-sericite-epidote
855-860	Chlorite-biotite-epidote
860-864	Chlorite-sericite-epidote-biotite
864-873	Sericite-epidote-biotite
873-876	Biotite-sericite-epidote
876-878	Chlorite-sericite-biotite
878-915	Biotite-sericite-epidote
915-920	Biotite-sericite-chlorite-epidote
920-930	Biotite-sericite-epidote
930-933	Chlorite-sericite-epidote-biotite
933-1018	Chlorite-epidote
1018-1026	Chlorite-sericite-epidote
1026-1027	Chlorite-epidote
1027-1036	Biotite-chlorite-epidote
1036-1060	Chlorite-sericite-epidote-biotite
1060-1072	Biotite-chlorite-epidote
1072-1075	Chlorite-epidote
1075-1085	Chlorite-epidote-biotite
1085-1100	Chlorite-biotite
1100-1104	Sericite-biotite-epidote

TABLE 1. (Cont'd)

<u>Footage</u>	<u>Mineral Assemblage</u>
1104-1113	Biotite-chlorite-sericite-epidote
1113-1130	Chlorite-biotite
1130-1160	Biotite
1160-1165	Biotite-chlorite-epidote
1165-1178	Chlorite-sericite-epidote-biotite
1178-1190	Biotite-sericite-chlorite
1190-1223	Biotite-sericite
1223-1226	Biotite-chlorite
1226-1250	Biotite-sericite-epidote
1250-1256-	Biotite-epidote
1256-1270	Biotite-chlorite-sericite
1270-1277	Biotite-sericite-epidote
1277-1295	Biotite-sericite-chlorite
1295-1299	Biotite

The interval with the highest total sulfide content observed was from 1030 to 1072 feet.

At 1146 feet a new mineral was first noted which has been tentatively identified as niccolite (NiAs), based on its association with pyrrhotite and its observed physical characteristics. It was found in assay reports that the interval containing the mineral had ten times the concentration of Ni than the general nickel background concentration of less than 10 ppm.

Framboids of pyrrhotite were also observed in the drill core from deeper portions of the hole. These later-formed concentrations of iron sulfide occur on fracture and foliation surfaces and imply a biogenic input for the reduction of sulfur and sulfide crystallization. Another unusual morphological feature of some of the sulfide minerals includes euhedral and subhedral pyrite cubes 1/8 to 1/4 inches on an edge disseminated in a meta-agglomeratic unit from 951 to 1006 feet. These are probably also late in the paragenetic sequence.

Oxide native copper is by far the major secondary oxide mineral, with traces of cuprite present only at one point. The copper is present in small amounts on fracture and foliation surfaces near the top of the drillhole. Higher concentrations are localized at 260 and 340 feet, and probably marks the position of earlier water tables. The last evidence of native copper is at about 500 feet.

#### STRUCTURAL FEATURES

Down DDH B-2 there were not so many top and bottom indicators as were observed in DDH B-1, possibly a result of the slightly higher metamorphic grade. However, two features were noted indicating structural complexity at Bromide.

Very small (about 1-2" amplitude, 1" wavelength) scale overturned, isoclinal drag folds were observed at 101 feet and 339 feet. These microfolds have an orientation similar to that proposed from evidence in DDH B-1. Since minor structures often times reflect regional trends, these small folds may aid in the understanding of the Bromide area.

In sawed sections of the core, in addition to the original bedding plane foliation, it became apparent that a second, less visible, preferred orientation of platy minerals exists. The lineation of biotite, and occasionally chlorite, that does not coincide with bedding surfaces suggests a later, minor folding event, possibly associated with the intrusion of igneous materials nearby.

A graphic explanation of this feature is shown in cartoon form in Figure 1. The interval of concern is between 422 and 425 feet. Because the vast majority of the core below this level is split rather than sawed, continuous monitoring of changes in this L2 lineation could not be accomplished. However, again at 1181 feet L1 and L2 were observed.

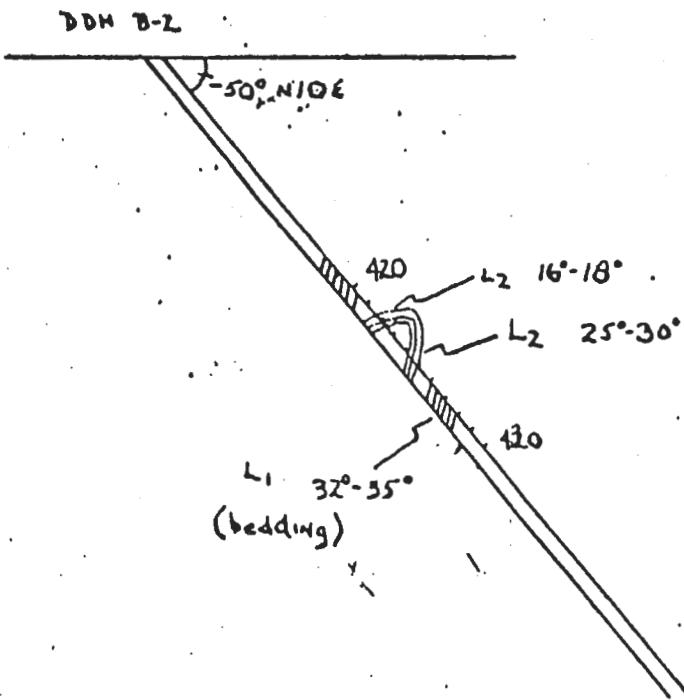


Figure 1.

— L<sub>1</sub>  
.... L<sub>2</sub>

R.W.S.  
18 JUNE 1978

U.S.BORAX DIAMOND DRILL LOG

## EXPLORATION - Detailed Alteration Log

## HOLE SURVEY

### *Footage*

## Bearing

### Inclination

Property Bromide

Hole No. B-2

Location Rio Arriba, N.M.

Total Depth 1299 ft.

Date Started \_\_\_\_\_

Collar Elevation ~97

Date Completed \_\_\_\_\_

Incl. & Bearing  $-50^\circ$ , N

Logged By RWS

Core Size NX

Foothills		Int. Rea.	Alteration #						Mineralization			Copper Ft.	Litho Lith	Geological Description		Assays
From	To		%	CHL	SCM	EP	DIA	CAL	T.S.	Ore	Mn.	Gangue		Rock Type, Structure, Etc.		
66															Becomes more "calc-like" in appearance	
66	20			45 ↓ 40 ↓ 15 ↓ 10 ↓ 10 ↓ 2 ↓ 4 ↓ 5 ↓ 5 ↓ 5 ↓ 10 ↓ 4											interlayered porphyritic and fine-grained equigranular volcanic flow material dominantly chlorite-epidote w/ calcite replacing plagi w/ epidote, limonitic veins, a few lg calcite pods - most are concordant w/ hydrothermal chlorite halo, minor scincite, a few gfg eyes	
86															- plumes become v. elongated - at 69' metamorphic Biotite, finely disseminated within chlorite-silicate-epidote schist picking up in abundance at depth, 1 small xenolith at 71' and becoming agglomeratic for about 2 foot interval	
86	88			35 ↓ 25 ↓ 5 ↓ 15								Hematite		then back to porphyritic mafic w/ cal. stringers - CLASTS are usually scincite rich - along calcite stringers - chrysotile, ad biotite increases (78'), Cu' assoc w/ porphy mafic usually w/ magnetite within agglomeratic clasts most plagi is completely cal-op replaced, very hematite-rich at ~ 85' and becomes agglomeratic		
88	99A			40 ↓ 15 ↓ 10 ↓ 10 ↓ 5 ↓ 45 ↓ 10 ↓ 7 ↓ 15										FAULT ZONE QUARTZ-CALCIITE-LIMONITE, HIGH URANIUM AND HYDROTHERMAL CHLORITE w/ EPIDOTE		
														finely laminated chlorite-mafic-ep schist w/ pelitic flow structures, interlayering of porphyritic and equigranular units ~ 1" thick, solution voids where pyrite has been oxidized, a few gfg eyes concordant calcite stringers, locally Biotite is fairly coarse, significant hydrothermal chlorite-chlorite after plagi, after scincite laminae		

Footage	Int.	Rea.	Alteration *					Mineralization			Arapahoe Log	General Description			Assays
			%	CHL	SER	EP	Bio	cal	T.S.	Ore Min.		Gangue	Fl.	Rock Type, Structure, Etc.	
99.4	102	2.6		50	5	15	3	5				mag octahedra and DUST		V.faint contact, porphyritic CHLOR-Ser- mag-calcite SCHIST, Limonite partially after Pyrite along foliation, minor interfoliate thrust faulting (WRT core the faults) Dip 40°-50° ∵ low-angle faults w/ horizontal Some re-katolized CHLORITE-hydrothermal. Locally high concentrations of SERICITE → Isoclinal-microfolds, truncated by inter faulting at 101'	
102	105	3		30	45	5	TR	3				v2% euhedral magnetite			
105	112.5	7.5		20	30	2	-	4				5% mag. nettle			
				↓	↓	↓									
				10	55	2									
				↓	↓	↓									
				20	35	1		8							
112.5	124.2	11.7		40	25	-	1	3				magnetite-dust		SERICITE-CHLORITE-magnetite-calcite SCHIST Locally rich in either SERICITE or CHLORITE RICH, Some metamorphic RIBBETTE going to vermiculite and Fe-OMOC after pyrite (?), Becomes more CHLORITE RICH w/ depth	
				↓											
				55	15	-									
				↓	↓										
				15	75	-	TR	5		Nat Cu <sup>+</sup>		coarser mag			
				15	45	-	2	5							
124.2	143	1		5	55	-	15	1-2				dis. fine mag		MULT QTS, Zoned CHLORITE, SERICITE, Hematite veining, ZONEO PLAGIOCLASE → POSSIBLY an ASHBLOW, Becomes more SERICITE w/ depth picks up pink Biotite (?) pods above w/ calcite near lower contact (111') w/ lenses of br. green chalc. and becomes more QTS near top, also magnetite lenses and generally coarser overall.	
				↓	↓	↓	↓	↓							
				5	35	1-2	30	1-2		tr Nat Cu					
138.5	139														
139															









Footage	Inch Rec.	Alteration *					Mineralization			Graphitic Log		General Description		Assays
		% CHL	See	EP	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	T.S.	Ore Min.	Gangue	Ft.				
From	To											Rock Type, Structure, Etc.		
259	263	4	15 ↓ 5	5 ↓ 45°	-	2 ↓ 2	1 ↓ 4	TR Cu <sup>+</sup> ASSOC w/ CHLORITE	minor mag-hem	abo		meta-sedimentite - ARKOSIC? - PINK FELDSP		
263	288	25	25 ↓ 20	10 ↓ 20	-	-	3 ↓ 3	TR Cu <sup>+</sup>	rounded mag-hem ~5%	270		grey WACCE (?) meta-schist, more chlorite and Limonite material, dominantly Qtz, Chalcocite, Sericitic, magnetite, plagioclase occasional pure chlorite lenses, magnetite Ring around rounded Qtz grain - replaced No real preferred orientation to grains, local Zones w/mixed Limonite & truncated grains Slow bedding → inverted, also graded Bedding supports this, rare Qtz pods at 275° hydrothermal chlorite & coarse Sericitic, often forming rims around pltg at 282° rare BiO mixed w/ Sericitic and chlorite becomes retrogressed, truncated Bedding, Right-side-up - evidence of 2 deformational episodes, one prior to deposition of omphacite unit, Biotite around hematite grains, small Qtz-ser pods with high magnetite, near lower contact Chlorite increases		
288	300	12	40 ↓ 60 ↓ 20	10 ↓ 5 ↓ 35	5 ↓ 5 ↓ 1	-	5 ↓ 3 ↓ 3	TR Py, Bn	magnetite ~2%	280		meta-andesite porphyry - CHALCOCITE RICH w/ASSOC pods of cal-Hydrothermal Chalc pltg is locally epioritized; becomes slightly brecciated w/ chalcocite matrix beds frags together - NOT much Porphyrobl., w/trace ASSOC SULFIDES → crackled zone?? high sericitic and biotite Chalc & ep become more pellitive, once out of pellitive chalc, no more "cracking" and sericitic zones, pltg is fresher MOST OF CHALCITE is hydrothermal cal... Sheathes CAL veins, get tetrahedrite at ~296°		







-see also Great

plot at -



Footage	Int.	Rec.	Alteration *					Mineralization			Graphitic Log	General Description			Assays	
			Chk	Ser	EP	Bio	cal	T.S.	Ore Min.	Gangue		Rock Type, Structure, Etc.				
From	To	%														
508.0	523.5											CONTINUE WITH SAME META-AGGLOMERATE UNIT - Qtz-SERICITE-BIORNITE-Hy. DEATHENED CHLORITE, APPEARS TO HAVE SERICITIZED STRETCHED PUMICEOUS (?) FRAGS, DISS. CALCIUM BIO IS DISS & UNALTERED IN BOTH CHLOR-ICH AND SER-ICH ZONES, MOST OF QTZ IS RE-KRISTALLIZED, SULFIDES ASSOC w/ BIOTITE MINOR EPIDOTE COMES IN AT 522', THIN LENSES LICH IN MAGNETITE, CHLORITE CONTENT INCREASES AT 527' AS DOES BIOTITE, MOSTLY ASSOC w/ CHLORITE, BY 520 CHLOR INCREASES AGAIN, ZONES OF OXIDATION - LIMONITE STAINED WITH CALCITE BEING, QTZ VERS				
523.5	525	1.5	5	60	-	12	+			MAG DUST	510					
					↓	↓	↓			after oxidized to hem or limonite						
			25	35	-	17	3				520					
			↓	↓	↓	↓	↓									
			10	55	-	12	8									
525	525		25	40	-	5	3			N54° mag-hem		High magnetite chlor-ser volcano-clastic (conglomerate ?) most chlor is hydrothermal				
525	561.5	365	10	65	-	3	2			Item after mag euhedral mag tetrahedra & octahedra		SEB INPUT				
			20	60	tr	4	2				530	BACK TO META-AGGLOMERATE w/ VENGET CONTROLLED BIOTITE, MINOR CHLOR IN A QTZ-CHL ROCK				
			↓	↓	↓	↓	↓					AFEW STRETCHED PEBBLES LOCALIZED MUCH. BIO				
			10	55	3	12	3					PICKS UP A LITTLE EPIDOTE AT 529' w/ CONCOMITANT INC IN CHLOR; LOCAL INC.				
					↓	↓	↓					IN BIOTITE, EPIDOTE INCREASING; LENSES OF SERICITIZED ASH w/ DISS BIOTITE. CLASTS INCREASE				
												IN SIZE AT 536' (f graded bouldy, then ↑)				
												WITH CONTINUED SER, BIO, EPIDOTE → MUDROCK				
			5	30	3	35	2					ROCK → RHODOLITE BRECCIA (?), BIO INCREASES				
			↓	↓	↓	↓	↓					SIGNIFICANTLY AT 540'; BECOMES A BIOTITE-				
			1	55	-	20	2					PIERCITE-QTZ ROCK WITH MINOR CALCIATE				
					↓	↓	↓					ONLY TRACES OF CHLORITE; FROM 542' TO 545'				
												INTERLAYING OF FELSIC ASH UNITS WITH				
			20	15	1-2	40	3					BIOTITE-CHLORITE AGGLOMERATE. ASH IS SERICITE				
			↓	↓	↓	↓	↓					WITH QTZ AND RETIC PLATE (?) AND DISSEMINATED				
			5	40	2	10	2					MAGNETITE, WITH BIOTITE VEINING (maybe a				
					↓	↓	↓					RHODOLITE?); HAS A BIOTITE-CHLORITE-RICH				
											ZONE NEAR BASE OF ASH, THEN GOES BACK TO					
			↓	60	-	4	1					SER-BIOTITE-QTZ AGGLOMERATE, A FEW SMALL				
												CALCIATE POOLS w/ CHLOR RIMS w/ A LITTLE EPIDOTE				
												DISSEMINATED AROUND BIOTITE GRAINS, ALSO EPIDOTE				
												REPLACING RETIC PLATE PHENES, A FEW ZONES				
												OF STRAW-COLORED SERICITE (PUMICE-ASSOC) w/ ONLY BIOTITE				
												VEINLETTS, BECOMES FINE GRAINED AGGLOMERATE,				
												BLACK ARE STRETCHED AND SINDLER, REKRISTALLIZED.				



Footage from To	Rec. %	Alteration *					Mineralization			Graphitic Log		General Description		Assays	
		CHL	Ser	Ep	Bio	Cal	T.S.	Ore Min.	Gangue	Ft.		Rock Type, Structure, Etc.			
13.5	621	7.5	5	-	5	25	2		tr py			altered diorite dike w/ chlorite borders — fine-grained, pheneritic, epidote veins, gtz-plag-magnetite, minor chlorite as trace py becomes coarser toward interior of dike, n/s inc in chl-ser, ep, mag; no preferred orientation — post metamorphic dike, Pyrite is widely disseminated			
21	625.5	44.6	tr ↓ 30 ↓ 1 ↓ 15 ↓ 1-2 ↓	↓ 20 ↓ 1 ↓ 1 ↓	microveins of py, ep	Diss euhedral mag ~ 8%				620		fine-grained meta-andelite, gtz-ser-chlor, epidote; significant magnetite; at 627' 6" gtz vein, local brecciation w/ biotite filling also magnetite, small gt eyes, short intervals of biotite schist, cut by cal-ep stringers, py assoc w/ cal rods rimmed by chlor-ep., also ep after plagi, @ ~637 gtz content increases to ~50% of rock, localized clusters of pyrite w/ cal-ep, at 642' about 6" of <u>andelite conglomerate</u> , Biotite veinlets sometimes as an envelope around cal-gtz stringers, another agglomerate unit 6" thick at 652', w/ clasts being gtz-ser rich & matrix of gtz-biot-sar, possibly some stretched porphyroblasts(?) clasts that are replaced by gtz, py assoc w/ gtz-ser Biotite is v. fine grained, and less abundant, most epidote confined to stringers, also coarser biotite is structurally (stringer, veinlet) controlled, trace amounts of pink Kspar			
			Z ↓ 30 ↓ - 20 ↓ 1 ↓	↓ 2 ↓ 1 ↓ 1 ↓	↓ 95 ↓ 1 ↓ 2 ↓				630						
			↓ 20 ↓ 8 ↓ 30 ↓ 1 ↓	↓ 1 ↓ 1 ↓ 1 ↓	↓ 20 ↓ 5 ↓ 20 ↓ 1 ↓				640						
			↓ 30 ↓ + 15 ↓ 2 ↓	↓ 1 ↓ 1 ↓ 1 ↓	↓ 30 ↓ 7 ↓ 5 ↓ 2 ↓	ass ep			650						
						tr py			660						
							mag. diat			670					
655	679.8		- 20 ↓ 15 ↓ 30 ↓ 5 ↓	↓ 10 ↓ 8 ↓ 35 ↓ 2 ↓	ass g (ba)				680		porphyritic andelite, v. small epidote-cal-plag phenos in a bio-sar-gtz; lge pads of gtz, biotite replacing mafic phenos pseudomorphically, w/ coarser biotite assoc w/ cal rods, veinlets, phenos are splotchy, isocubeled-anhedral, and become more abundant (~30% of rock) w/ stockwork of cal-ep veinlets, inc in biotite & coarsening at 675', and then plants thin out toward lower contact, and bio slightly increased in matrix & in veinlets				
			TR ↓ 20 ↓ 5 ↓ 40 ↓ 3 ↓		tr py										























Footage	Int.	Rec.	Alteration *					Mineralization			Graphic Log	General Description		Assays	
			From	To	%	CHL	SER	EP	BIO	CAL		T.S.	Ore Min.	Gangue	
1223.5	1226.5	3				25	-	-	+5	.12			Diss Po		
126.5	128.5	12	TR	40	-	20	3						veinlet Po		
			2	40	+	20	3								
			1	25	2	35	2							dis. mag	
			7	15	5	45							Po, NiAs		
128.5	1243.5	6	3	15	10	35	6						Po-veinlet asso		
			4	30	6	20	5						1/2% at 25		
248.5	1250	6.5	TR	15	6	30	3						tarnished Po		
			3	5	2	40	4						Diss & veinlet ~3% of rock		
			TR	2	4	95	6								
250	1256	6	3	10	TR	40	5						veinlet Po		
			15	5	-	30	3						—		

